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Context and issues - Coupling of acoustic and elastic wave equations

Goals

- Accurate simulation of seismo-acoustic waves through **heterogeneous domains with complex geometries**
- Treatment of realistic cases of interest
 - **High computational costs**

Issues

- High-order precision needed to accurately capture waves
 - **Hybrid discontinuous methods (HDG/HHO)**

- Acoustic wave equations:**
$$\begin{cases} \rho_F \partial_t \mathbf{v}^F(t) + \nabla p(t) = \mathbf{0} \\ \frac{1}{\kappa} \partial_t p(t) + \nabla \cdot \mathbf{v}^F(t) = g(t) \end{cases}$$
- Elastic wave equations:**
$$\begin{cases} \partial_t \boldsymbol{\varepsilon}(t) - \nabla^s \mathbf{v}^S(t) = \mathbf{0} \\ \rho_S \partial_t \mathbf{v}^S(t) - \nabla \cdot (\mathbf{C} : \boldsymbol{\varepsilon}(t)) = \mathbf{f}(t) \end{cases}$$
- Coupling conditions:**
$$\begin{cases} \llbracket \mathbf{v}(t) \cdot \mathbf{n}_\Gamma \rrbracket = 0 \\ (\mathbf{C} : \boldsymbol{\varepsilon}(t)) \cdot \mathbf{n}_\Gamma = p(t) \mathbf{n}_\Gamma \end{cases}$$

Application of HHO method to seismo-acoustic coupling

Approximation spaces:

- Acoustic domain:** $\mathbf{v}^F \rightarrow \text{dG}(k)$, $p \rightarrow \text{HHO}(k', k)$
- Elastic domain:** $\boldsymbol{\varepsilon} \rightarrow \text{dG}(k)$, $\mathbf{v}^S \rightarrow \text{HHO}(k', k)$

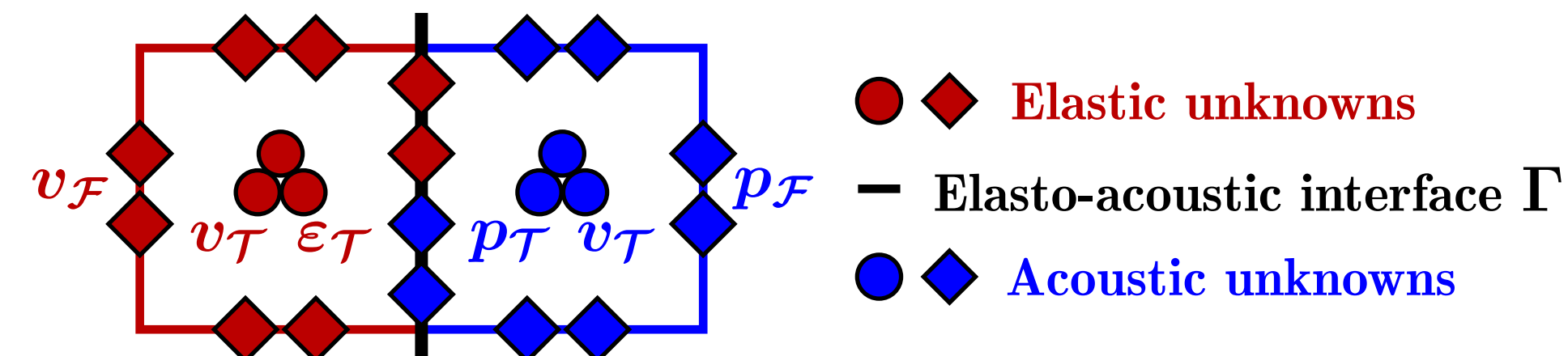


Fig. 1: Elasto-acoustic unknowns with a mixed-order discretization ($k' = k + 1 = 2$)

Algebraic realization:

$$\begin{bmatrix} M_{\mathcal{T}\mathcal{T}}^V & 0 & 0 & 0 & 0 & 0 \\ 0 & M_{\mathcal{T}\mathcal{T}}^F & 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & M_{\mathcal{T}\mathcal{T}}^S & 0 & 0 \\ 0 & 0 & 0 & 0 & M_{\mathcal{T}\mathcal{T}}^S & 0 \\ 0 & 0 & 0 & 0 & 0 & 0 \end{bmatrix} \begin{bmatrix} \partial_t \mathbf{V}_{\mathcal{T}}^F \\ \partial_t P_{\mathcal{T}} \\ \partial_t P_{\mathcal{F}} \\ \partial_t S_{\mathcal{T}} \\ \partial_t V_{\mathcal{T}} \\ \partial_t V_{\mathcal{F}} \end{bmatrix} + \begin{bmatrix} 0 & -G_{\mathcal{T}} & -G_{\mathcal{F}} & 0 & 0 & 0 \\ G_{\mathcal{T}}^\dagger & \Sigma_{\mathcal{T}\mathcal{T}}^F & \Sigma_{\mathcal{T}\mathcal{F}}^F & 0 & 0 & 0 \\ G_{\mathcal{F}}^\dagger & \Sigma_{\mathcal{F}\mathcal{T}}^F & \Sigma_{\mathcal{F}\mathcal{F}}^F & 0 & 0 & C_{\mathcal{T}} \\ 0 & 0 & 0 & 0 & -E_{\mathcal{T}} & -E_{\mathcal{F}} \\ 0 & 0 & 0 & E_{\mathcal{T}}^\dagger & \Sigma_{\mathcal{T}\mathcal{T}}^S & \Sigma_{\mathcal{T}\mathcal{F}}^S \\ 0 & 0 & -C_{\mathcal{T}}^\dagger & E_{\mathcal{F}}^\dagger & \Sigma_{\mathcal{F}\mathcal{T}}^S & \Sigma_{\mathcal{F}\mathcal{F}}^S \end{bmatrix} \begin{bmatrix} \mathbf{V}_{\mathcal{T}}^F \\ P_{\mathcal{T}} \\ P_{\mathcal{F}} \\ S_{\mathcal{T}} \\ \mathbf{V}_{\mathcal{T}}^S \\ \mathbf{V}_{\mathcal{F}}^S \end{bmatrix} = \begin{bmatrix} 0 \\ G_{\mathcal{T}} \\ 0 \\ 0 \\ \mathbf{F}_{\mathcal{T}} \\ 0 \end{bmatrix}$$

Mechanical energy of the scheme: $\mathcal{E}_h(t) := \mathcal{E}_h^S(t) + \mathcal{E}_h^F(t)$ with

$$\mathcal{E}_h^F(t) := \frac{1}{2} \|\mathbf{v}_{\mathcal{T}}^F(t)\|_{L^2(\rho_F; \Omega_F)}^2 + \frac{1}{2} \|p_{\mathcal{T}}(t)\|_{L^2(\frac{1}{\kappa}; \Omega_F)}^2, \quad \mathcal{E}_h^S(t) := \frac{1}{2} \|\mathbf{v}_{\mathcal{T}}^S(t)\|_{L^2(\rho_S; \Omega_S)}^2 + \frac{1}{2} \|\boldsymbol{\varepsilon}_{\mathcal{T}}(t)\|_{L^2(\mathbf{C}; \Omega_S)}^2$$

Semi-discrete energy conservation of the scheme:

$$\mathcal{E}_h(t) = \mathcal{E}_h(0) + \int_0^t \left[(\mathbf{f}(\alpha), \mathbf{v}_{\mathcal{T}}^S(\alpha))_{L^2(\Omega_S)} + (g(\alpha), p_{\mathcal{T}}(\alpha))_{L^2(\Omega_F)} - s_h^S(\hat{\mathbf{v}}_h^S(\alpha), \hat{\mathbf{v}}_h^S(\alpha)) - s_h^F(\hat{p}_h(\alpha), \hat{p}_h(\alpha)) \right] d\alpha$$

Propagation of an acoustic (water) pulse into an elastic medium (granite)

Computational parameters:

- Computational domain:**
 - Water on the upper side
 - Granite on the lower side
- Mixed-order discretization:** $k' = k + 1 = 3$
- Time integration scheme:** SDIRK(3,4)
- Time step:** $dt = 0, 1 \times 2^{-9}$
- Homogeneous Dirichlet conditions**
- IC:** velocity Ricker wave in the acoustic medium: $\mathbf{v}_0^F(x, y) := \theta \exp\left(-\pi^2 \frac{r^2}{\lambda^2}\right) (x - x_c, y - y_c)^T$

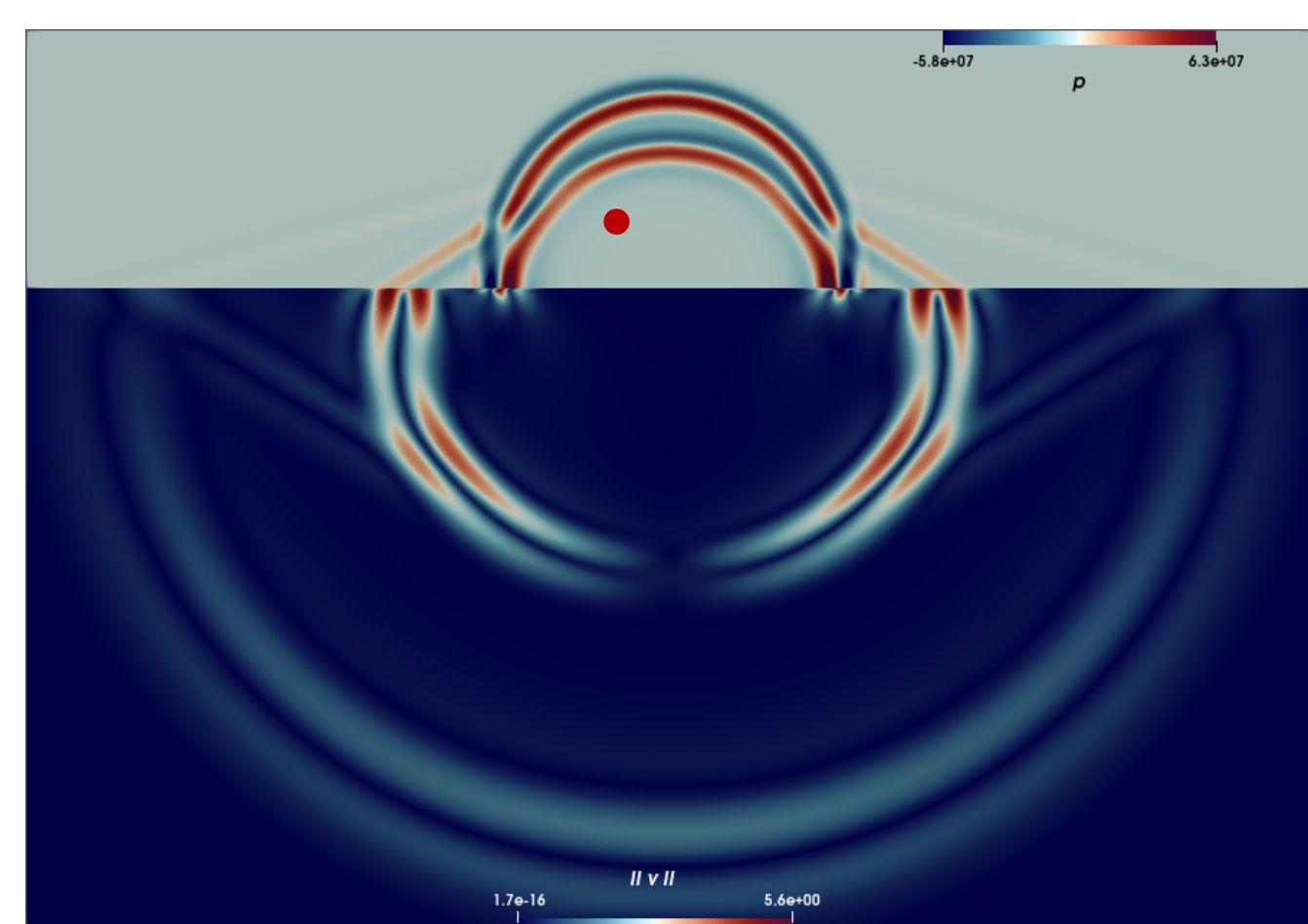


Fig. 5: Left panel: Distribution of acoustic pressure and elastic velocity norm, at $t = 0, 4375$ s. Right panel: Pressure as a function of time at a sensor in the water (coarse mesh)

General principles of HHO method

HHO is a finite element method similar to the Hybrid Discontinuous Galerkin method (HDG)

Design of HHO method

- Degrees of freedom:**
 - Polynomial unknowns located in the cells (degree k') and on the faces (degree k): $\hat{\mathbf{u}}_h := (\mathbf{u}_{\mathcal{T}}, \mathbf{u}_{\mathcal{F}})$
 - Equal-order discretization: $k' = k$
 - Mixed-order discretization: $k' = k + 1$
- Operators:**
 - Gradient reconstruction operator: $\nabla \mathbf{u} \rightarrow \mathbf{G}(\hat{\mathbf{u}}_h)$,
 - Stabilization operator: $\mathbf{s}(\hat{\mathbf{u}}_h, \hat{\mathbf{w}}_h)$
 - Penalty at element level to enforce weakly the matching of the trace of the cell unknown with the local face unknowns

Advantages over classical finite element methods

- Mesh flexibility:**
 - Complex geometries
 - Unstructured and polyhedral meshes
 - Local mesh refinement
- Local conservativity**
- Optimal error estimates (for smooth solutions)**
- Attractive computational costs:**
 - Global problem couples only face dofs
 - Cell dofs recovered by local post-processing

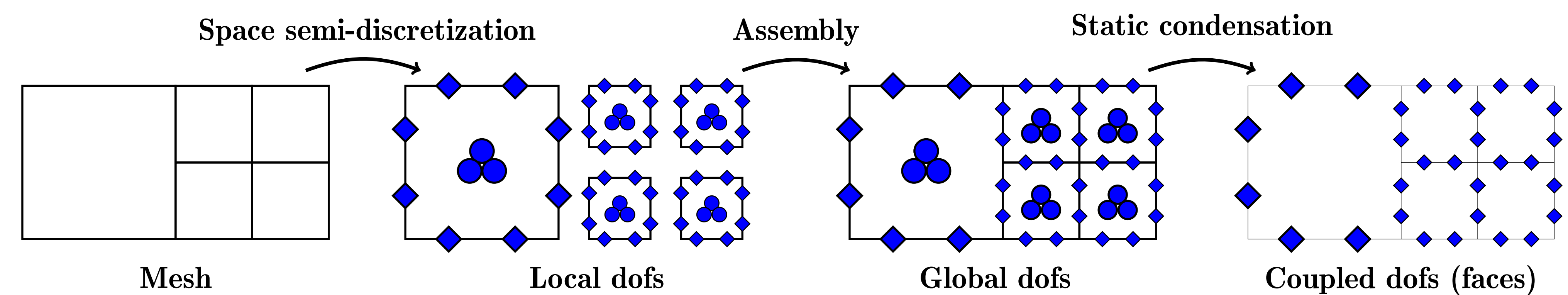


Fig. 2: Static condensation procedure

Verification on academic test cases

Convergence rates on sinusoidal solutions:

- $\mathcal{O}(h^{k+1})$ in H^1 -norm
- $\mathcal{O}(h^{k+2})$ in L^2 -norm (superconvergence)

Energy conservation:

- $\rho_S = \rho_F = 1$, $c_P^S = c_P^F = \sqrt{3}$, $c_S^S = 1$
- **IC:** velocity Ricker wave in acoustic medium:

$$\mathbf{v}_0^F(x, y) := \theta \exp\left(-\pi^2 \frac{r^2}{\lambda^2}\right) (x - x_c, y - y_c)^T$$

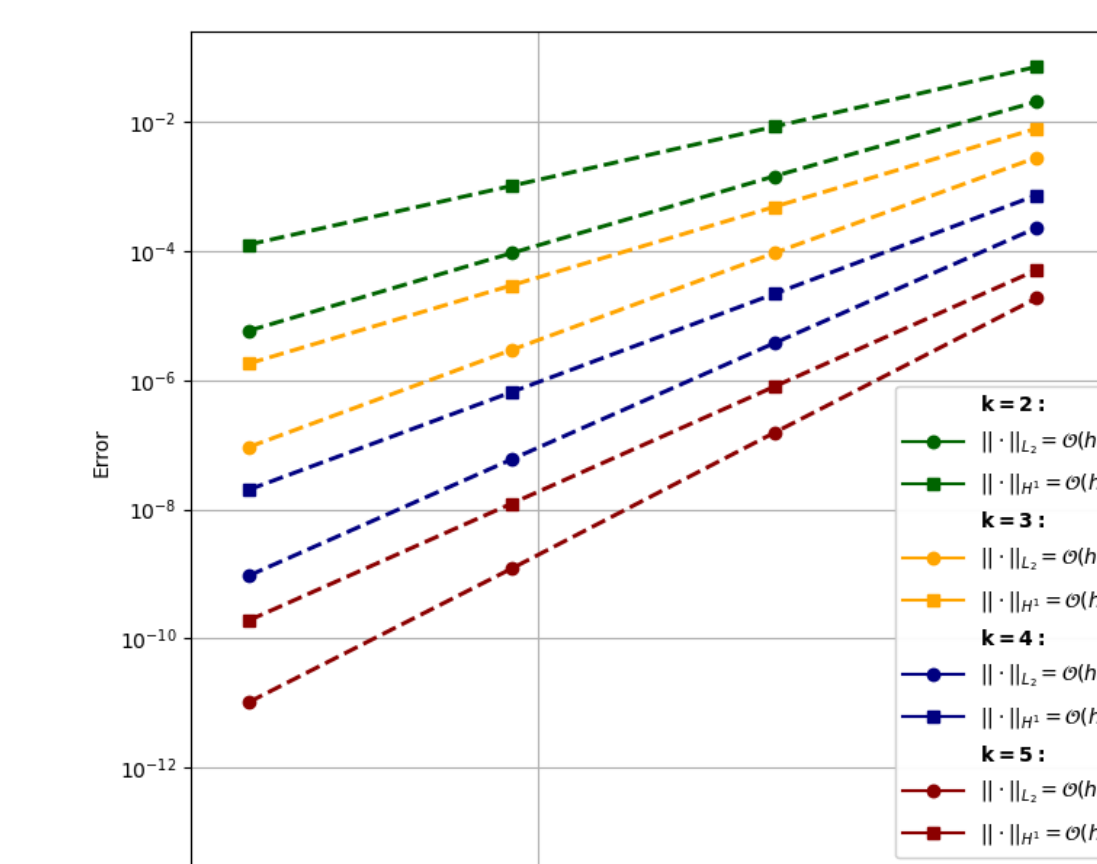


Fig. 3: Errors as a function of the mesh size with $\Delta t = 0, 1 \times 2^{-5}$

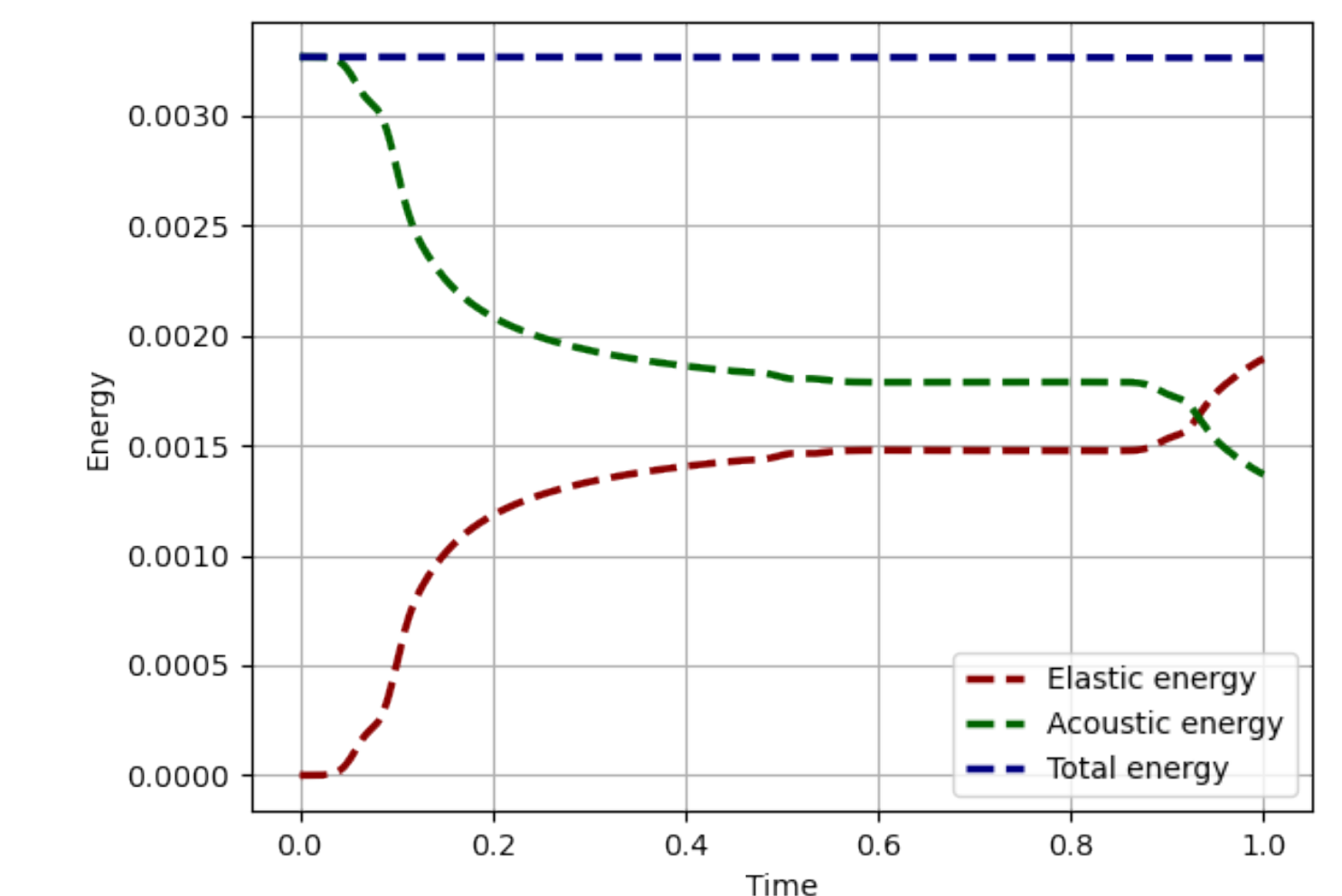


Fig. 4: Repartition of elastic and acoustic energy

Propagation of an elastic pulse in a sedimentary basin and atmosphere

- Acoustic region:** $\rho = 1, 292 \text{ kg.m}^3$, $c_p = 340 \text{ m.s}^{-1}$
- Sedimentary region:** $\rho = 1200 \text{ kg.m}^3$, $c_p = 3400 \text{ m.s}^{-1}$, $c_s = 1400 \text{ m.s}^{-1}$
- Bedrock region:** $\rho = 5350 \text{ kg.m}^3$, $c_p = 3090 \text{ m.s}^{-1}$, $c_s = 2570 \text{ m.s}^{-1}$

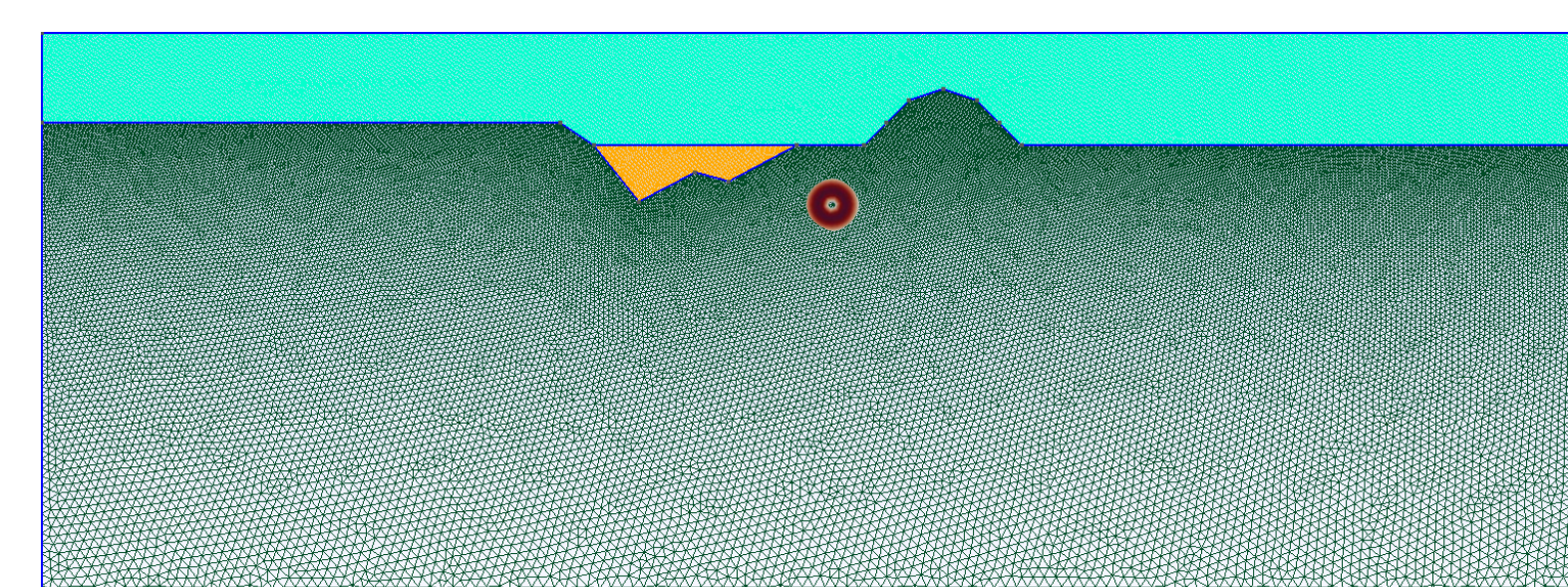


Fig. 6: Mesh of sedimentary basin and location of initial pulse

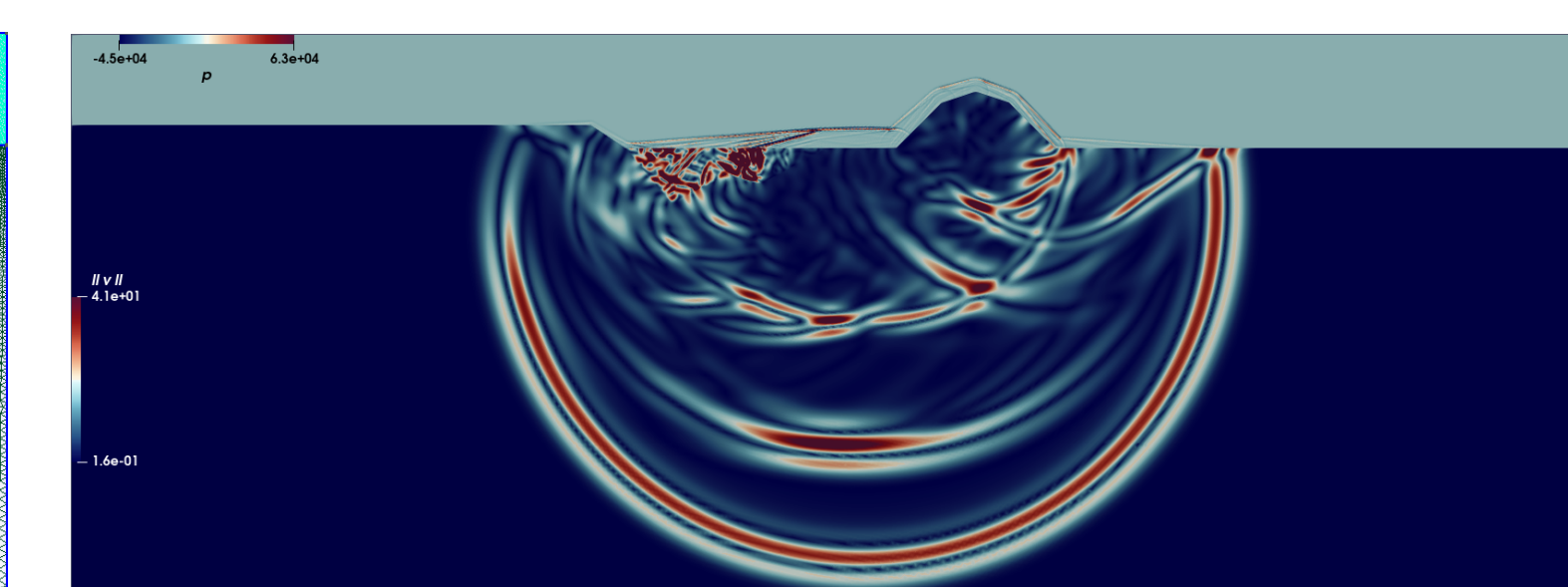


Fig. 7: Propagation through the sedimentary basin

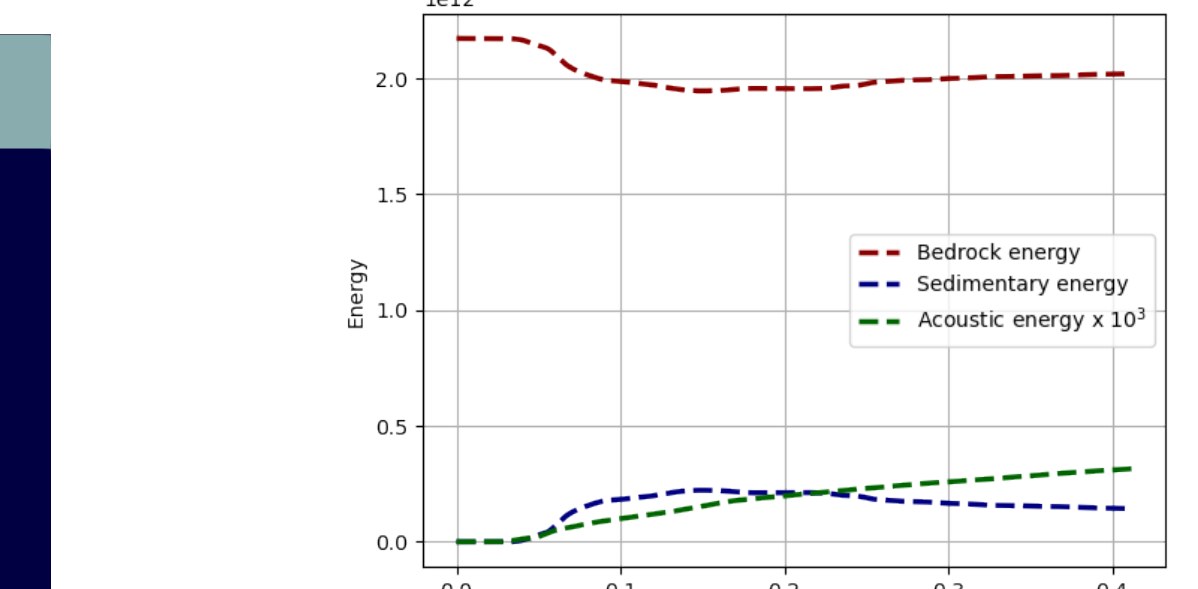


Fig. 8: Repartition of elastic and acoustic energy

Energy transfer enhancement above the sedimentary basin

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